

Effect of Sub-pixel Clouds and Aerosols on the retrieval of Trace Gases

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Effect of Clouds on the Retrieval of BL Trace Gas column

$$\hat{N} = \frac{SCD}{AMF}; \quad AMF = (1 - w) * AMF_{clear} + w * AMF_{cloud}$$

where, w is the "cloud fraction"

For clouds above polluted BL: $AMF_{cloud} \approx 0$

$$\hat{N} \approx \frac{SCD}{(1 - w) AMF_{clear}}$$

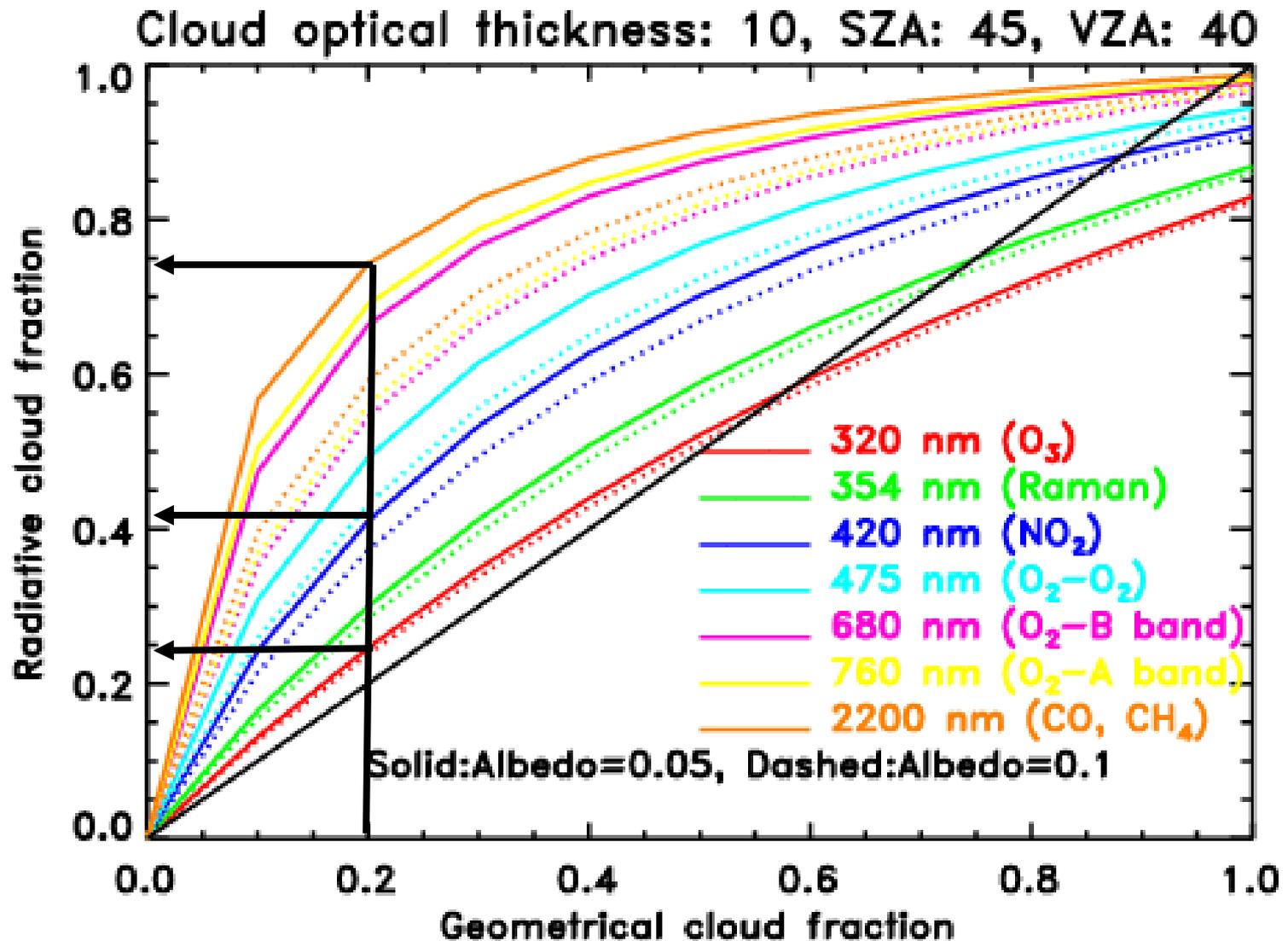
- Since clouds are usually much brighter than the surrounding atmosphere w can be several times larger than the geometrical cloud fraction.
- 10% error in w can produce 10-100% error in the estimated vert. column depending upon w .
- Cloud height is less important unless clouds are inside the BL.

What is “Cloud Fraction”?

- For trace gas measurement at reflected λ s it is the fraction of the measured radiation that comes from clouds (not including below cloud radiance that propagates thru clouds).
 - The OMI team calls it Cloud Radiance Fraction or Radiative Cloud Fraction (w).
- Formally, $w = f_g * \rho_c / \rho_m$
where f_g is the geometrical cloud fraction, ρ_c is the reflectance of the “bare” cloud, and ρ_m is the TOA reflectance corrected for atm absorption.

Cloud Radiance Fraction varies with λ

Rayleigh scattering effect on w



A Simple Method to Calculate w

Using Mie theory it can be shown that w varies linearly with ρ_{clear} / ρ_m

Applying boundary conditions :

$$w = \left(\frac{\rho_m - \rho_{clear}}{\rho_m^* - \rho_{clear}} \right) \left(\frac{\rho_m^*}{\rho_m} \right)$$

where, ρ_m^* is the "critical" TOA reflectance above which $w = 1$

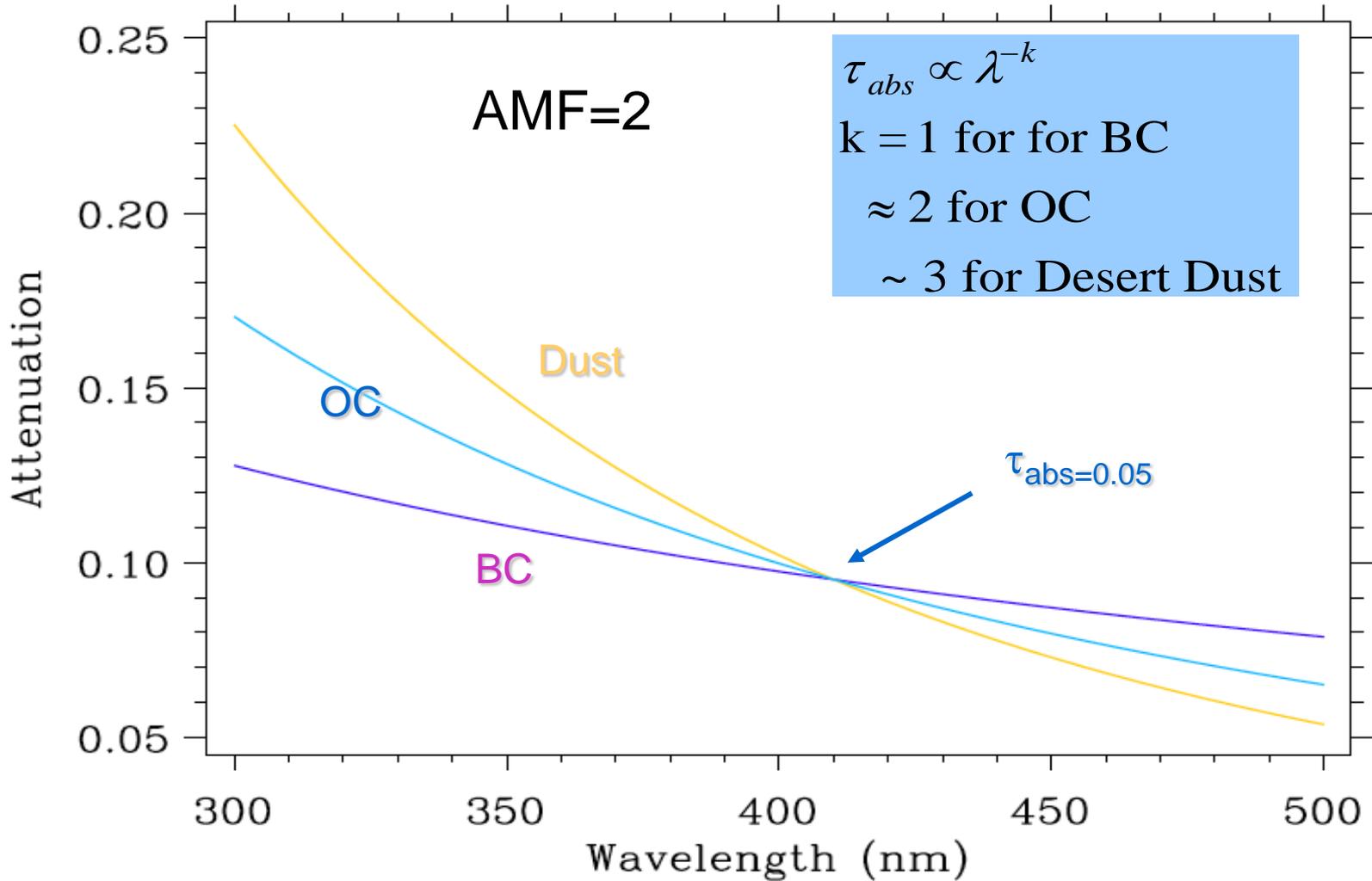
Since $\rho_m^* \gg \rho_{clear}$, w is insensitive to ρ_m^*

In the TOMS and OMI algorithms ρ_m^* is calculated assuming an 80% reflecting Lambertian cloud. The results agree well with Mie theory.

Correction for Aerosols

- OMI method of calculating w accounts for non-absorbing aerosols of any kind.
- If aerosols have absorption, ρ_{clear} is overestimated, causing w to be underestimated.
 - the effect is worse in the UV than in the visible
 - the effect is worse for elevated aerosols
 - the effect is worse over brighter surfaces
- UV absorbing aerosol index (UV-AAI) can be used to develop a correction.

Reduction in w by Aerosol Absorption

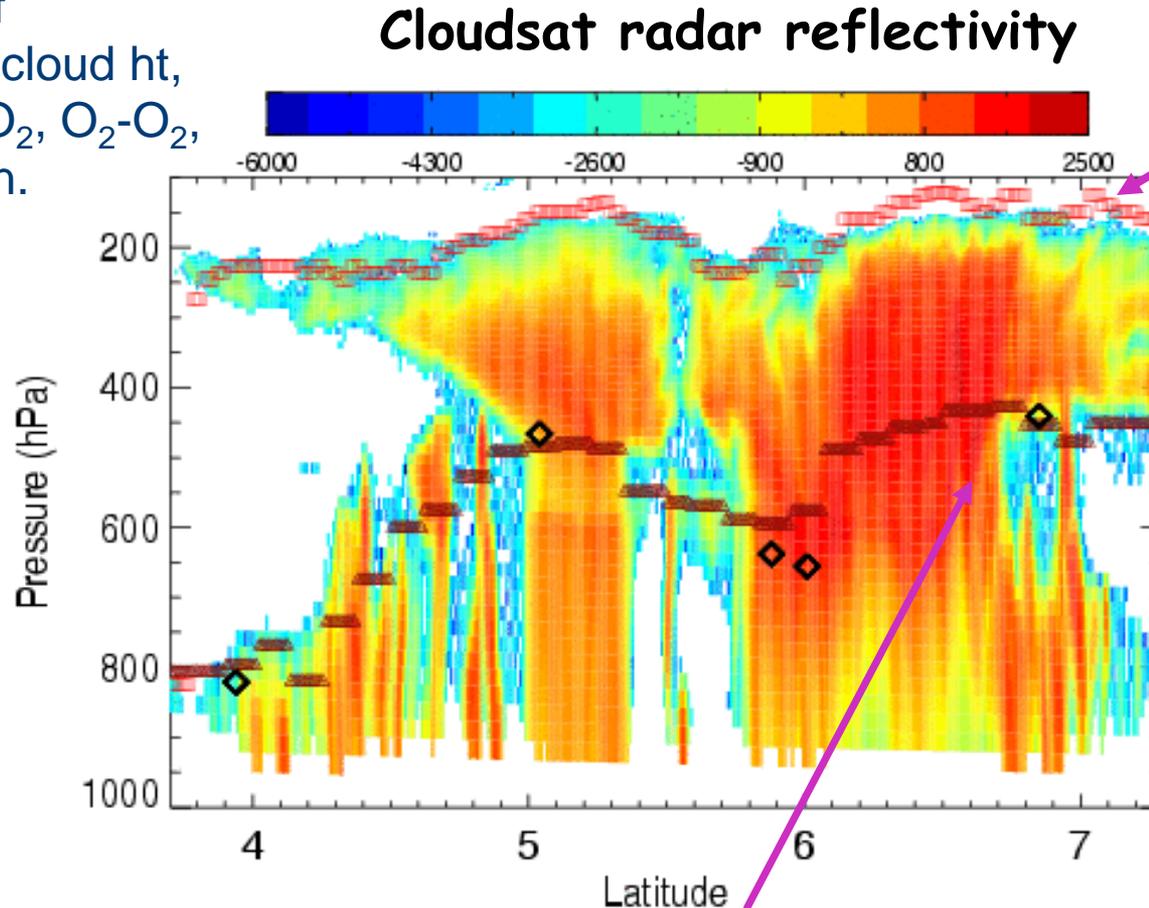


What About Cloud height?

- At reflected wavelengths sunlight penetrates into clouds and passes through most types of clouds.
 - Hence trace gas absorption is affected by cloud vertical structure. Cloud/aerosol mixed scenes can produce complicated effects.
- Concept of Optical Centroid Pressure (OCP) has been developed to account for this complexity.
 - The concept is accurate only if the trace gas is well mixed between the scattering layers, but is ALWAYS more accurate than assuming cloud-top pressure to estimate AMF.

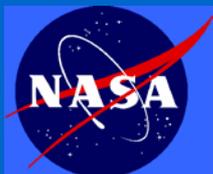
Multi-phase/Multi-layer Cloud Effects

There are several methods of estimating cloud ht, e.g., TIR, O_2 , O_2-O_2 , and Raman.



MODIS cloud-top press is insensitive to cloud vertical structure

Cloud Optical Centroid press calculated using OMI-measured Rot Raman Scattering is sensitive to cloud vert structure
(ref : Vasilkov *et al.*, JGR, '08)



Cloud OCP Retrieval Methods

- OMI uses Rotational Raman scattering @350 nm and O₂-O₂ absorption @477 nm to estimate OCP.
 - Both methods agree well with Cloudsat-derived OCP when $w > 0.5$.
 - Random and systematic errors in OCP increase as $1/w$ for small w .
 - Both methods overestimate OCP when w is small.

Retrieved BL trace gas columns are probably most accurate when $0.3 < w < 0.7$

Summary

- BL trace gas retrieval requires accurate knowledge of Cloud Radiance Fraction (w).
 - A simple method to calculate w , developed independently at NASA and KNMI, works very well, but correction for absorbing aerosols needs to be developed.
- To estimate the AMF it is far better to use OCP than the cloud-top pressure.
 - The two OMI methods for estimating OCP underestimate the OCP of clouds and aerosols for $w < 0.5$.
- Aerosol and cloud mixed scenes present a significant challenge.
 - Currently, there is no good method to handle them.